# Longevity Estimate of the Sidoarjo Mud Eruption 'Lusi'

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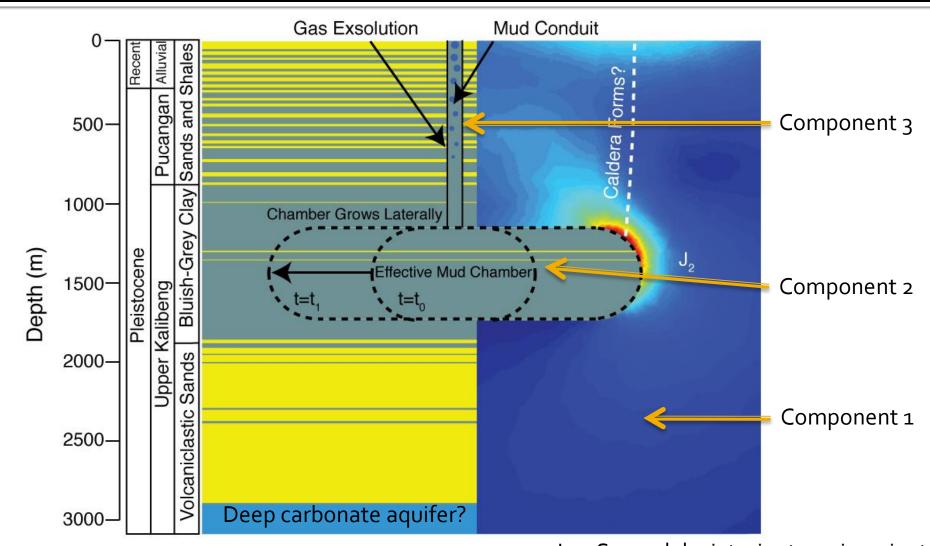
## When will the eruption end?

- We need to understand how the mud source functions.
- We develop a model of the mud source and plumbing system.
- Certain unknown parameters affect the model outcome. We take the Monte Carlo approach to explore the sensitivity of the outcome to each of these parameters.
- We make statistical predictions of longevity.

# **Key Assumptions**

- The mud comes from the upper Kalibeng, between 1200-1800 meters depth
  - Microfossils, Temperature, Clay mineralogy (Mazzini et al.2007), Geodesy (Fukushima et al. 2009)
- Water comes from upper Kalibeng after initial stage
  - Consistent with 30% porosity (Tanikawa et al. 2010). If Davies et al. (2011) value of 10-13% is more realistic, a deep aquifer is needed to explain 30% water content at surface (Bayuaji et al. 2009). An additional water source may have played a role during the initial stage of the eruption.
- Mud exhibits a yielding rheology
  - Laboratory rheology experiments (Manga et al. 2009, Rudolph and Manga 2010), Field observations (Kopf et al. 2009), also inability of relief wells to penetrate mud source (Sawolo personal communication).

#### **Model - Schematic**



Stratigraphy from Mazzini et al. 2007

J2 – Second deviatoric stress invariant

### Model – Mud Reservoir

#### Free Surface

Momentum Equation

$$\nabla \cdot \underline{\sigma} = \underline{0}$$

Constitutive Equation  $\sigma_{ij} = \lambda \epsilon_{kk} \delta_{ij} + 2\mu \epsilon_{ij}$ .

$$\sigma_n|_{\partial S} = \Delta P_C$$

Mud Chamber

(Uniform Pressure)

Yielding Criterion:

Yielding occurs according to von Mises criterion. Chamber pressure tends to decrease as material erupts, but low hydraulic diffusivity of surroundings keep pore pressures high. When surroundings yield and become incorporated into chamber, chamber pressure increases.

The von Mises stress is a scalar measure of deviatoric stresses

Pressure-Volume relation:

$$\rho_0 V_{0,C} + \int_0^t \dot{M}(\tau) d\tau = V_C(P) / v_S(P)$$

P chamber pressure,

 $V_{C}(P)$  current deformed chamber volume

 $v_s(P)$  specific volume of (particles+liquid+gas)

Initial mass of system – (Cumulative mass removed) = Mass of material remaining in chamber

#### Model – Mud Conduit

$$P=P_{\text{atm}}$$

$$\frac{\partial}{\partial z} \left(\rho_{mix} u\right) = 0$$

$$\frac{\partial P}{\partial z} = \frac{-\rho_{mix} \left(g + \frac{8\mu u}{\rho_{mix} r^2}\right)}{1 - \frac{u^2}{c^2}}$$

$$c = \left(\frac{\partial P}{\partial \rho_{mix}}\right)_S^{1/2}$$

We model the conduit in 1D

Atmospheric pressure at top boundary

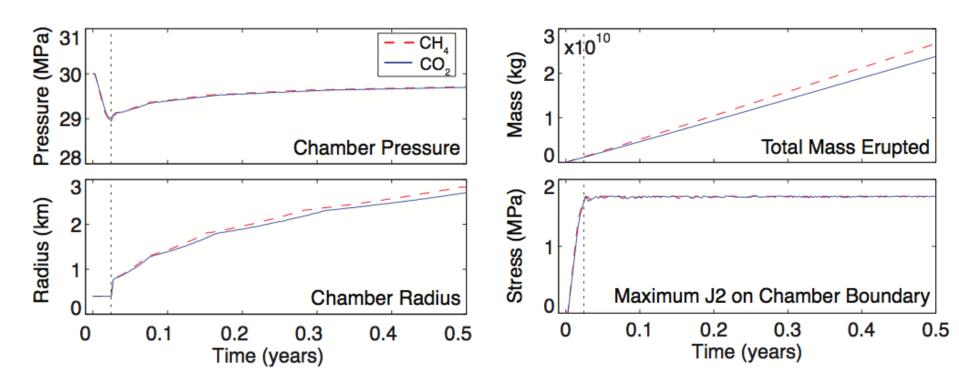
Chamber pressure at bottom boundary

Conservation of mass and momentum

Sound speed computed from equation of state

We use the CO<sub>2</sub>-CH<sub>4</sub>-H<sub>2</sub>O equation of state of Duan et al. 1992a,b.

### Results – Example



#### Key observations:

- 1) Yielding begins at dashed line
- 2) After yielding, chamber begins to expand
- 3) After yielding, stresses near chamber are held constant at the yield stress
- 4) Discharge is very nearly constant

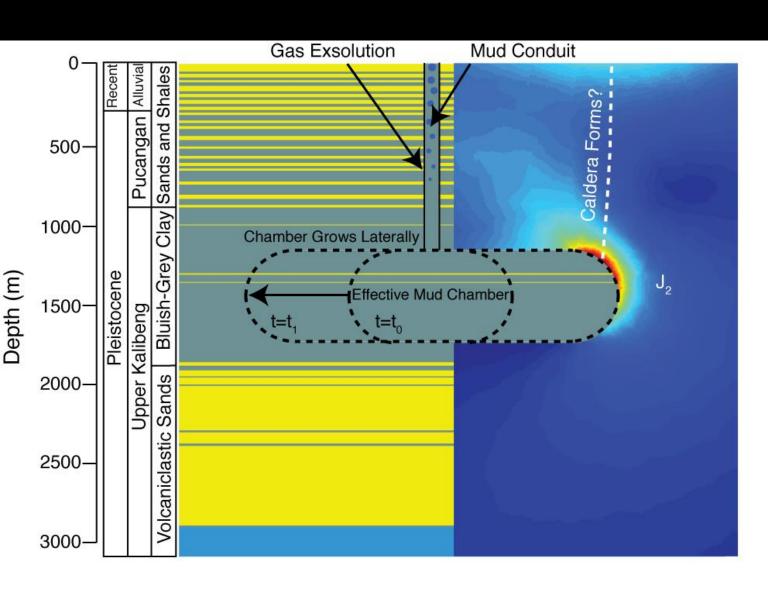
# **Model – Eruption Termination**

#### Two possibilities:

- 1) Chamber pressure becomes too small.
- 2) A caldera forms.

Eruption may not truly end when caldera forms, but our model is no longer applicable. We assume that a caldera forms when J<sub>2</sub> exceeds a critical value in the region between the tip of the chamber and the land surface.

#### Unknowns

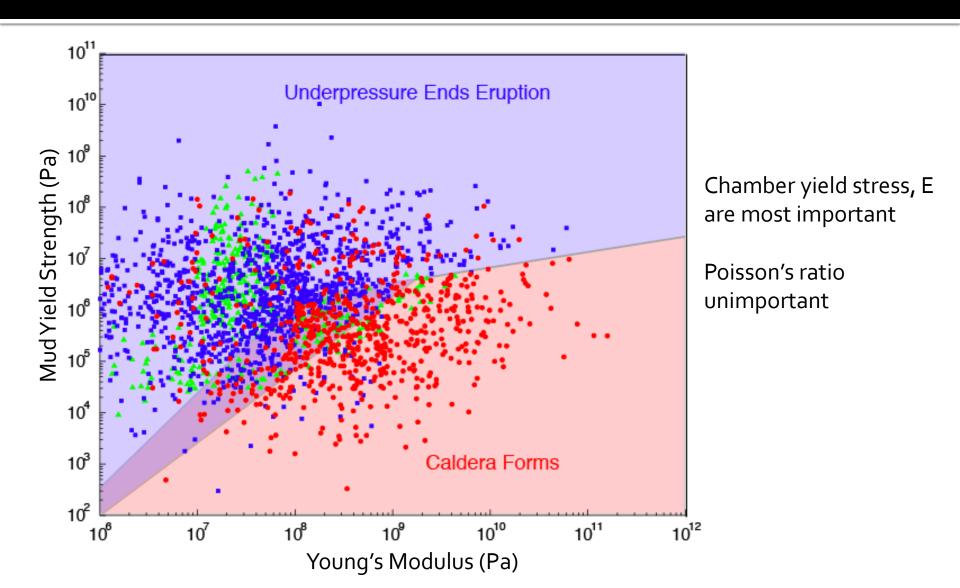


#### **Monte Carlo Simulations**

Parameter	Constant or Variable?	Range (1 standard deviation) or value
E (Young's modulus)	Variable	10 <sup>7</sup> -10 <sup>9</sup> Pa
ν (Poisson's ratio)	Variable	.0525
Mud Yield Strength	Variable	10 <sup>5</sup> -10 <sup>7</sup> Pa
Yield Strength for Caldera	Variable	10 <sup>6</sup> -10 <sup>8</sup> Pa
Water Content (by Volume)	Constant	30 %
Mud Viscosity	Constant*	10 <sup>4</sup> Pa-s
Conduit Radius	Constant*	1.4 m ( $CH_4$ ) or 8.5 m ( $CO_2$ )
Gas Concentration	Constant	1% CO <sub>2</sub> or 0.5% CH <sub>4</sub>

<sup>\*</sup>For a given gas composition and viscosity, we iteratively find a radius that satisfies observational constraints

#### Results – Sensitivity to Parameters



## Results – Longevity Predictions

#### Three Scenarios:

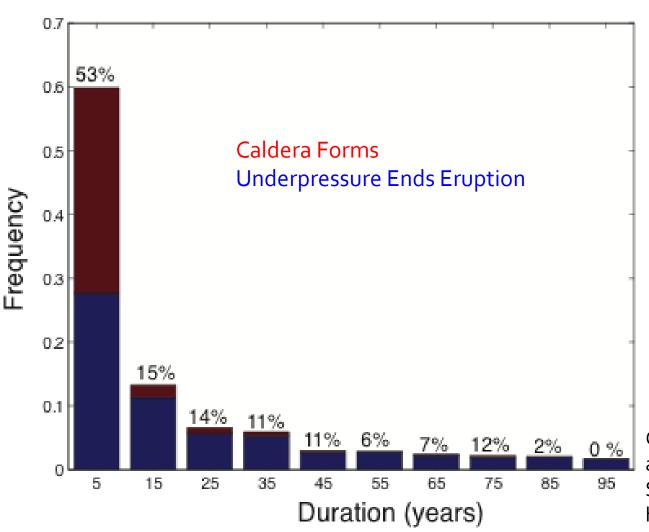
- 1) Gaussian variable distributions
- 2) Uniform variable distributions within 1 standard deviation
- 3) Uniform variable distributions within 2 standard deviations

Scenario:	1 - Gaussian	2 — Flat (1σ)	3 — Flat (2σ)
33%	21	27	14
50%	40	50	25
66%	84	>100	52

We assume here that the eruption lasts more than 5 years.

Note: Our predictions are somewhat longer than those of Davies et al. 2011

#### Results - Outcome



Longer eruptions are less likely to produce a caldera.

Complete results in manuscript to appear in Earth and Planetary Science Letters. Available at: http://seismo.berkeley.edu/~max

# How can we improve predictions?

- Geochemical constraints on:
  - 1) Source of water and relative contributions from deep and shallow sources.
  - 2) How the relative contributions have changed over time. We can perform analyses, but need samples.
- Better estimates of discharge as a function of time.
- Constraints on mud source rheology (yield strength). We can measure this, but need samples.
- Constraints on mud source geometry.